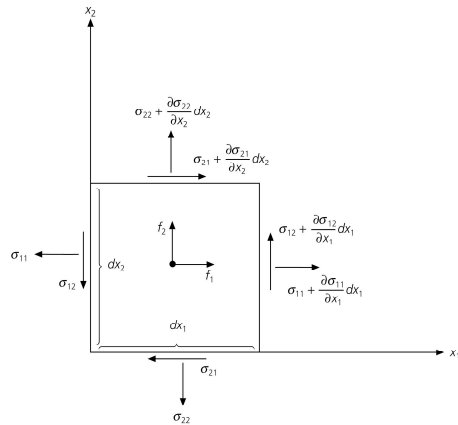


Seismic wave equation

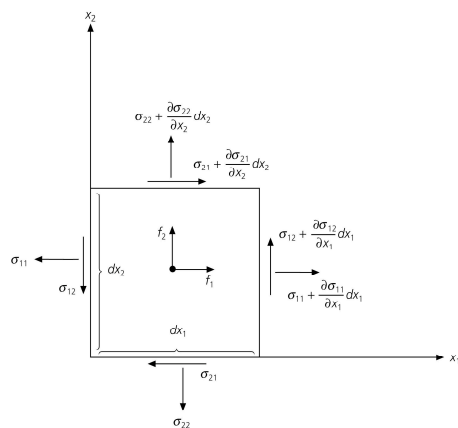


$$\text{mass} \times \text{acceleration} = \sum \text{forces}$$

$$\rho \frac{\partial^2 u_i}{\partial t^2} = \sigma_{ij,j} + f_i$$

$$\rho \frac{\partial^2 u_i}{\partial t^2} = c_{ijkl} u_{(k,l),j} + f_i$$

Seismic wave equation



$$\rho \frac{\partial^2 u_i}{\partial t^2} = c_{ijkl} u_{(k,l),j} + f_i$$

isotropic medium:

$$c_{ijkl} = \lambda \delta_{ij} \delta_{kl} + \mu (\delta_{il} \delta_{jk} + \delta_{ik} \delta_{jl})$$

Helmholtz decomposition:

$$\underline{u} = \underline{\nabla} \phi + \underline{\nabla} \times \underline{\psi}$$

$$\underline{\nabla} \times \underline{\nabla} \phi = 0$$

$$\underline{\nabla} \cdot \underline{\nabla} \times \underline{\psi} = 0$$

Seismic wave equation

P-wave

wave eq: $\alpha^2 \nabla^2 \phi - \frac{\partial^2 \phi}{\partial t^2} = -\frac{1}{\rho} F_p$

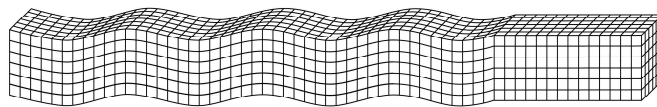
velocity: $\alpha = \sqrt{\frac{\lambda + 2\mu}{\rho}}$

S-wave

wave eq: $\beta^2 \nabla^2 \psi - \frac{\partial^2 \psi}{\partial t^2} = -\frac{1}{\rho} F_s$

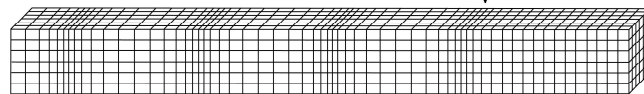
velocity: $\beta = \sqrt{\frac{\mu}{\rho}}$

S waves: ground motion is perpendicular to wave direction



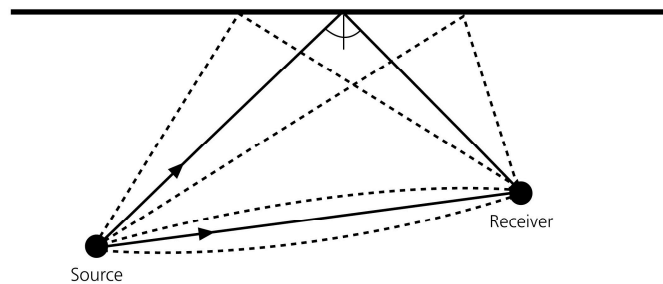
Direction of wave propagation

Onset of waves

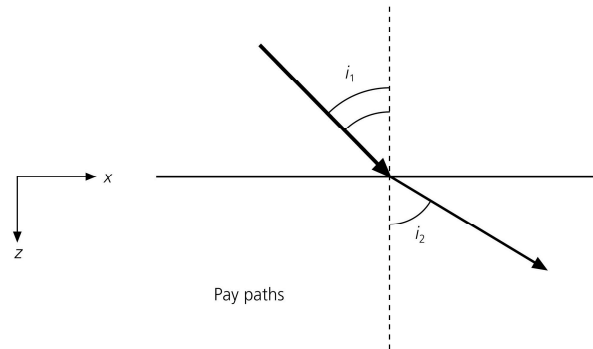


P waves: ground motion is parallel to wave direction

Ray theory

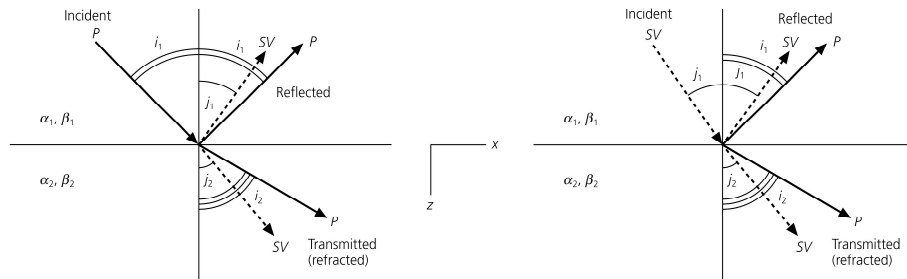


Snell's law



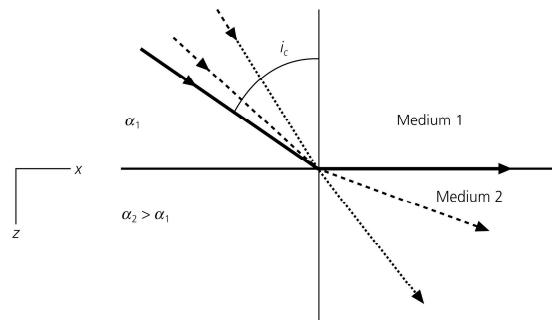
$$p = \frac{\sin i_1}{v_1} = \frac{\sin i_2}{v_2}$$

Snell's law



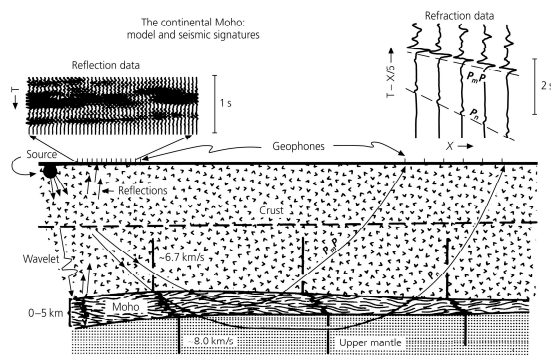
$$p = \frac{\sin i_1}{\alpha_1} = \frac{\sin j_1}{\beta_1} = \frac{\sin i_2}{\alpha_2} = \frac{\sin j_2}{\beta_2}$$

Head waves

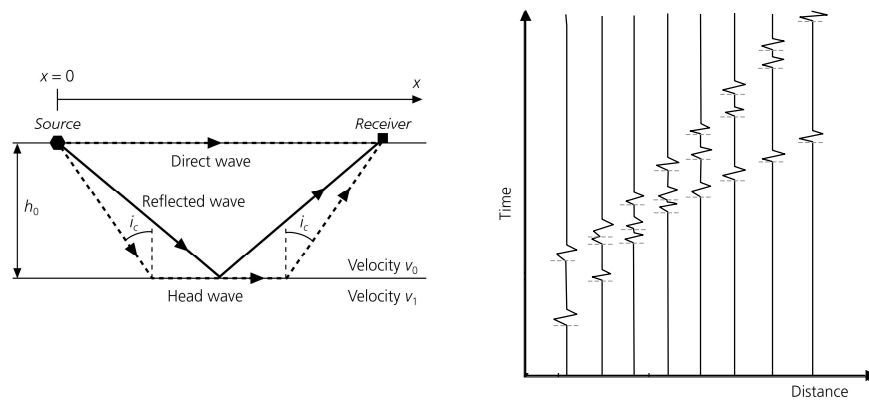


$$p_c = \frac{\sin i_c}{\alpha_1} = \frac{\sin(\pi/2)}{\alpha_2} = \frac{1}{\alpha_2}$$

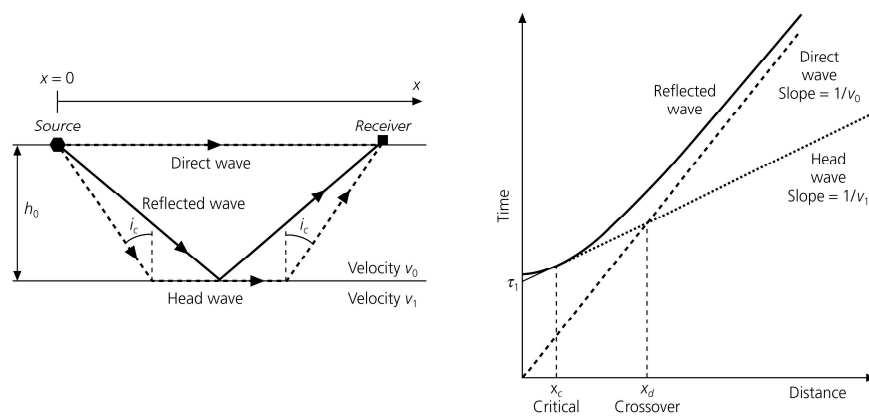
Seismic reflection and refraction



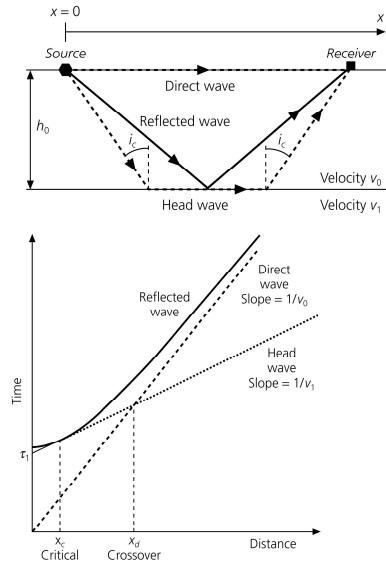
Seismic refraction



Seismic refraction



Seismic refraction



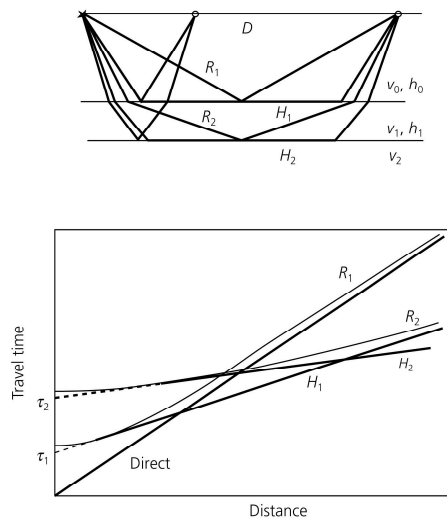
$$T_D(x) = \frac{x}{v_0}$$

$$T_R(x) = 2 \frac{\left(\frac{x^2}{4} + h_0^2\right)^{1/2}}{v_0} = px + 2\eta_0 h_0$$

$$T_H(x) = \frac{x}{v_1} + 2h_0 \left(\frac{1}{v_0^2} - \frac{1}{v_1^2} \right)^{1/2} = \frac{x}{v_1} + \tau_1$$

where vertical slowness is: $\eta_0 = \left(1/v_0^2 - p^2\right)^{1/2}$

Seismic refraction



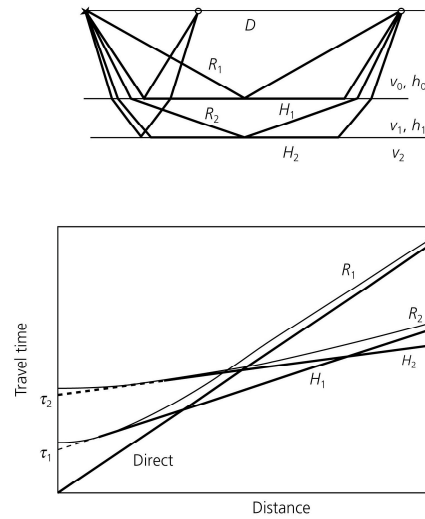
$$T_D(x) = \frac{x}{v_0}$$

$$T_{Rn}(x) = px + 2 \sum_{j=0}^{n-1} \eta_j h_j$$

$$T_{Hn}(x) = \frac{x}{v_n} + 2 \sum_{j=0}^{n-1} h_j \left(\frac{1}{v_j^2} - \frac{1}{v_n^2} \right)^{1/2} = \frac{x}{v_n} + \tau_n$$

$$h_{n-1} = \frac{\tau_n - 2 \sum_{j=0}^{n-2} h_j \left(\frac{1}{v_j^2} - \frac{1}{v_n^2} \right)^{1/2}}{2 \left(\frac{1}{v_{n-1}^2} - \frac{1}{v_n^2} \right)^{1/2}}$$

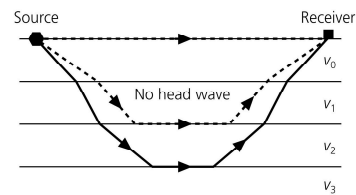
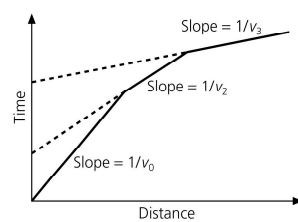
Seismic refraction



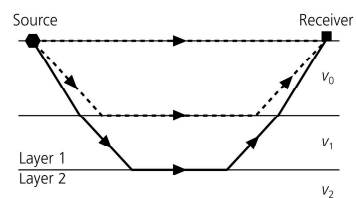
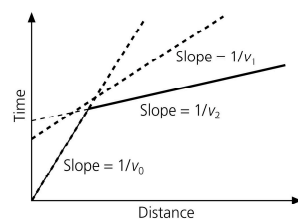
$$h_{n-1} = \frac{\tau_n - 2 \sum_{j=0}^{n-2} h_j \left(\frac{1}{v_j^2} - \frac{1}{v_n^2} \right)^{1/2}}{2 \left(\frac{1}{v_{n-1}^2} - \frac{1}{v_n^2} \right)^{1/2}}$$

Seismic refraction

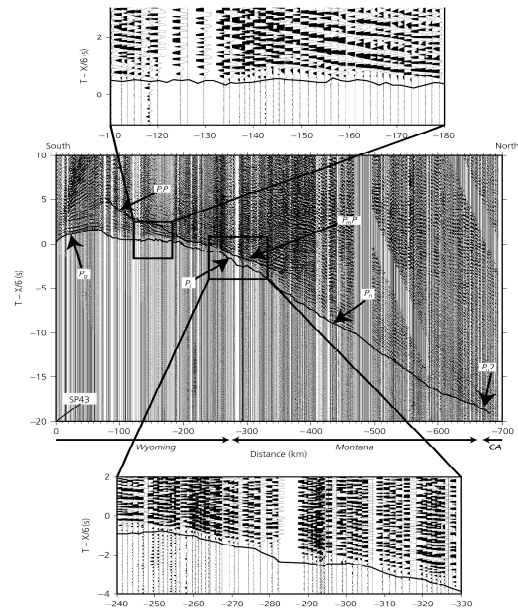
Low-velocity layer:



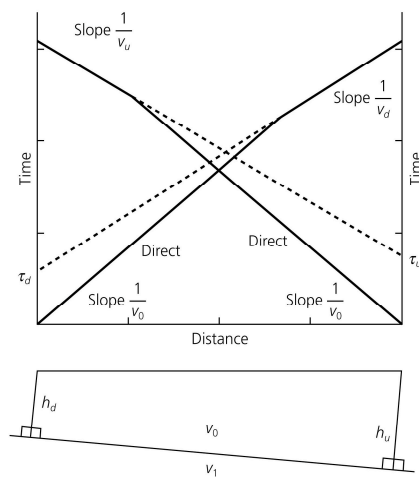
Blind zone:



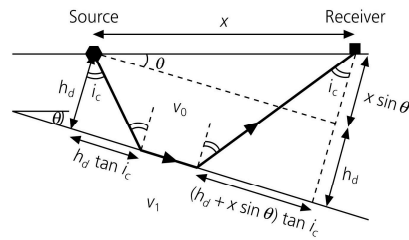
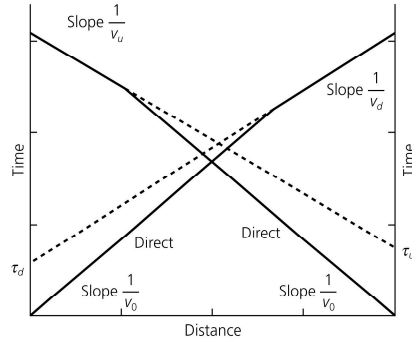
Seismic refraction



Seismic refraction



Seismic refraction



$$T_d(x) = \frac{x \cos \theta - (2h_d + x \sin \theta) \tan i_c}{v_1} + \frac{(2h_d + x \sin \theta)}{v_0 \cos i_c}$$

$$= \frac{x \sin(i_c + \theta)}{v_0} + \frac{2h_d \cos i_c}{v_0}$$

$$= \frac{x}{v_d} + \tau_d$$

$$T_u(x) = \frac{x \cos \theta - (2h_u - x \sin \theta) \tan i_c}{v_1} + \frac{(2h_u - x \sin \theta)}{v_0 \cos i_c}$$

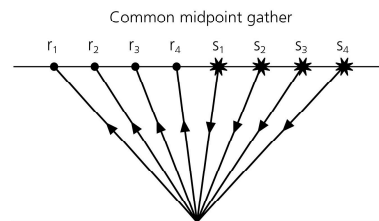
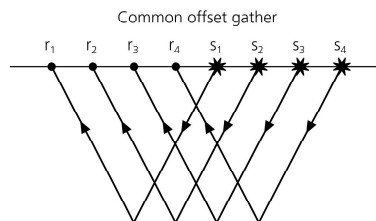
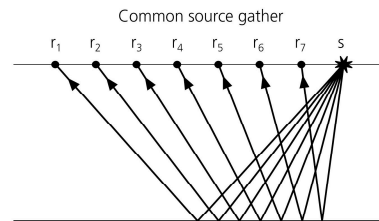
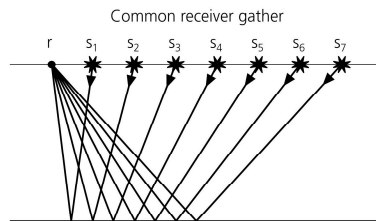
$$= \frac{x \sin(i_c - \theta)}{v_0} + \frac{2h_u \cos i_c}{v_0}$$

$$= \frac{x}{v_u} + \tau_u$$

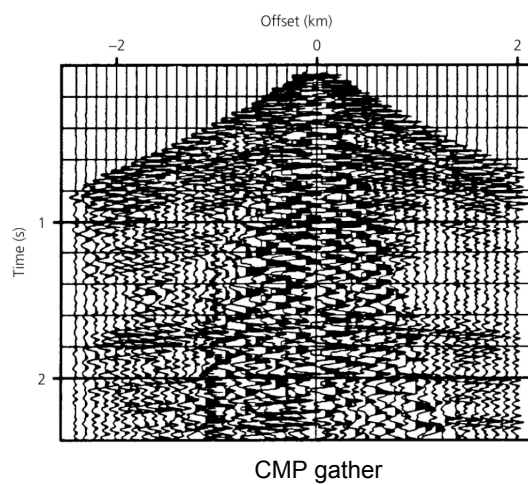
$$\theta = \frac{1}{2} \left(\sin^{-1} \frac{v_0}{v_d} - \sin^{-1} \frac{v_0}{v_u} \right)$$

$$i_c = \frac{1}{2} \left(\sin^{-1} \frac{v_0}{v_d} + \sin^{-1} \frac{v_0}{v_u} \right)$$

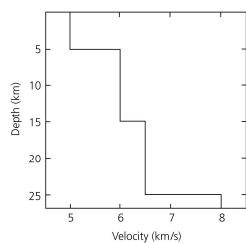
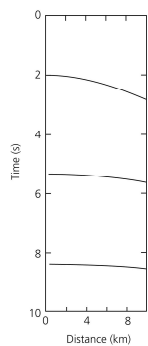
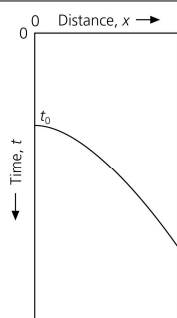
Seismic reflection



Seismic reflection



Seismic reflection



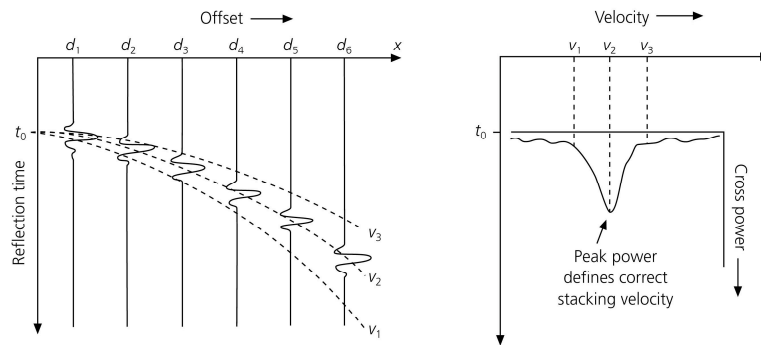
Normal moveout (NMO) correction:

$$\Delta T_{NMO}(x) = T_{Rn}(x) - t_n$$

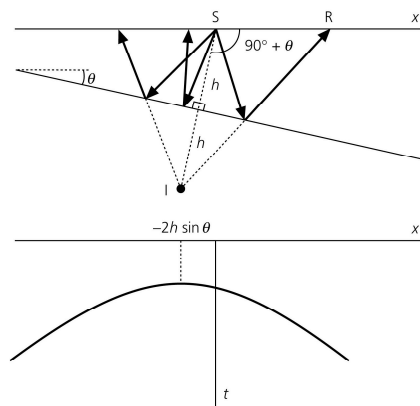
$$= px + 2 \sum_{j=0}^{n-1} \eta_j h_j - t_n$$

$$t_n = 2 \sum_{j=0}^n h_j / v_i$$

Seismic reflection



Seismic reflection



$$T_R(x) = \frac{(x^2 + 4h^2 + 4hx \sin \theta)}{v_0^2}$$

Seismic reflection

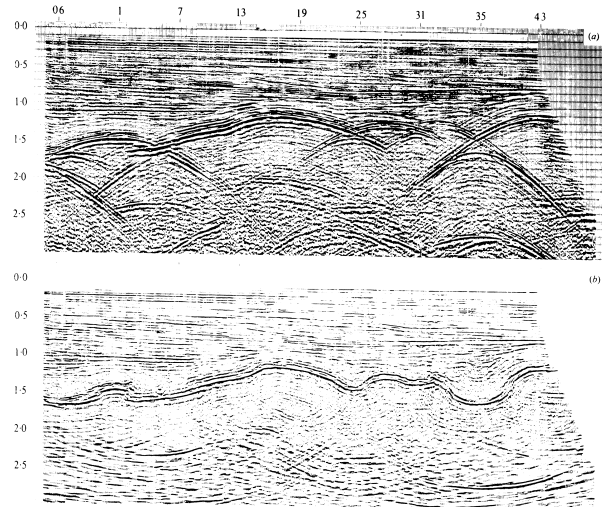


Fig. 9.57 Migration removes the confusion produced by multiple-branch reflections. The deeper data, which are probably multiples, out-of-plane diffractions, and other types of noise

are smeared out because they are migrated as if they were primary reflections. (Courtesy of AGIP) (a) Unmigrated section; (b) migrated.

Sheriff and Geldart, 1994